



Conference Proceedings of the 4<sup>th</sup> Asia Pacific Luminescence and Electron Spin Resonance Dating Conference  
Nov 23<sup>rd</sup>-25<sup>th</sup>, 2015, Adelaide, Australia

Guest Editor: Mathieu Duval

## THE ALPHA EFFECTIVENESS OF THE DATING ESR SIGNAL IN BARITE: POSSIBLE DEPENDENCE WITH AGE

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Received 14 February 2016

Accepted 19 October 2016

**Abstract:** The alpha effectiveness value (k-value) for the ESR (Electron Spin Resonance) signal due to  $\text{SO}_3^-$  in barite was revised by comparing the dose responses of the signal intensities to gamma rays and to 4 MeV  $\text{He}^+$  ion doses in natural sea-floor hydrothermal barite samples actually used for dating. Of the values obtained for a synthetic, a natural old, and a natural young samples, the one for the natural young sample is tentatively adopted, which is  $0.053 \pm 0.006$ , although further works are still necessary to establish this value.

**Keywords:** barite, hydrothermal activities, ESR dating, alpha effectiveness.

### 1. INTRODUCTION

Barite ( $\text{BaSO}_4$ ) is a mineral recently found to be practically useful for ESR dating (Okumura *et al.*, 2010; Takamasa *et al.*, 2013; Fujiwara *et al.*, 2015), especially for those occurring in submarine hydrothermal sulfide deposits, although Kasuya *et al.* (1991) first proposed that dating of barite is possible. Barite crystals are formed when submarine hydrothermal fluid containing Ba is mixed with sea water containing sulfate anions. As the hydrothermal fluid contains large amount of Ra, it incorporates in barite replacing Ba in the crystal lattice. In the hydrothermal sulfide deposits containing barite, most of the natural radiation is from  $^{226}\text{Ra}$ ,  $^{228}\text{Ra}$  and their daugh-

ter radioactive nuclei where other radioactive elements such as U, Th, and K in sulfide minerals give negligible contribution to the dose rates (Toyoda *et al.*, 2014). When estimating the dose rate to barite crystals, which is a necessary procedure for ESR dating, therefore, internal radiation, especially, that of alpha particles is significant, typically, 40 to 60% of total dose rates (Okumura *et al.*, 2010; Fujiwara *et al.*, 2015).

In the procedure of ESR dating, the age,  $T$ , is obtained by

$$T(\text{ka}) = \frac{\text{Equivalent dose, } D_E (\text{Gy})}{\text{Dose rate, } D (\text{mGy/y})} \quad (1.1)$$

when  $D$  is constant, or

$$D_E = \int_0^T D(t) dt \quad (1.2)$$

when  $D$  changes with time. The dose rate,  $D$ , is given by

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$$D = kD_{\alpha} + D_{\beta} + D_{\gamma} + D_{\text{cos}} \quad (1.3)$$

where the dose rate is the sum of doses of alpha, beta, gamma, and cosmic ray dose rates denoted as  $D_{\alpha}$ ,  $D_{\beta}$ ,  $D_{\gamma}$  and  $D_{\text{cos}}$ , respectively. It should be noted that, here, the  $D_E$  and  $D$  are the effective values, that are calibrated by the gamma rays. As alpha particles have a large LET (linear energy transfer), they produce high concentration of electron-hole pairs, hence, higher probability of recombination, *i.e.*, lower effectiveness of trapped unpaired electrons (ESR signal). The ratio,  $k$ , is the alpha effectiveness, which is the signal formation rate in the dose response by alpha particles relative to that by gamma rays, has to be used to correct this difference to estimate the “effective” dose rates.

Toyoda *et al.* (2012) investigated the alpha effectiveness for the ESR signal due to  $\text{SO}_3^-$  in a barite crystal formed on land, by comparing the dose responses of the signal for gamma irradiation and for  $\text{He}^+$  ion implantation with energy of 4MeV, which simulates the alpha particles. A value  $0.043 \pm 0.018$  was obtained for a sample from Morocco. However, the dose response for  $\text{He}^+$  ion dose is far from “good”. The experiments of  $\text{He}^+$  ion implantation need be repeated to determine the precise alpha effectiveness. In the present paper, we implanted the  $\text{He}^+$  ions accelerated by a Tandem accelerator into two series of barite samples, one synthetic and the other hydrothermal extracted from a sea-floor hydrothermal sulfide deposit. The alpha effectiveness of the  $\text{SO}_3^-$  is determined as the slope in the dose response to  $\text{He}^+$  ion implantation relative to that to gamma ray irradiation, denoted as

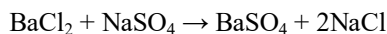
$$k = \frac{\text{Slope (He ion implantaion)}}{\text{Slope (Gamma ray irradiation)}} \quad (1.4)$$

## 2. EXPERIMENTAL

A natural barite sample was extracted from a hydrothermal deposit taken from the sea-floor at the Hatoma Knoll, and Iheya North Knoll of Okinawa Trough as listed in **Table 1** where they were dated by ESR to be 1480 to 6300 years. A block of each sulfide deposit sample was cut into pieces, and about 2.0 g was crushed. The samples were soaked in 12M hydrochloric acid, left for approximately 24 hours. Then, 13M nitric acid was add-

ed. Finally, after rinsing in distilled water, the sample was filtered and dried. Impurities were removed by handpicking. An X-ray diffraction study was made to confirm that the grains are pure barite. The samples were further crushed to powder.

Another barite was synthesized by mixing barium chloride ( $\text{BaCl}_2$ ) and sodium sulfate ( $\text{Na}_2\text{SO}_4$ ) solutions, as the reaction expressed by



Aqueous solutions of 300 ml with 0.32 mol/L of barium chloride and of 300 ml with 0.32 mol/L of sodium sulfate, are mixed together at room temperature to have barium sulfate precipitated.

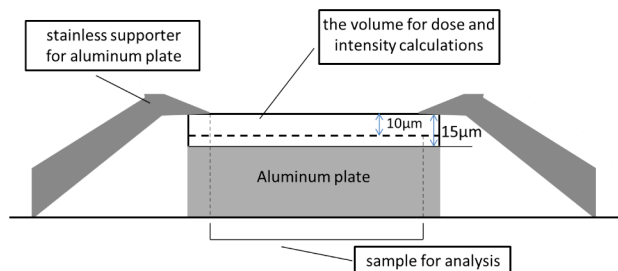
The above powder barite samples of two types (natural and synthetic) were deposited on aluminum plates in deionized water with an area of 2 cm<sup>2</sup> where the thickness of the barite was about 15  $\mu\text{m}$  (6.75 mg/cm<sup>2</sup>) while the range of 4 MeV  $\text{He}^+$  ion in barite is about 10  $\mu\text{m}$  (4.5 mg/cm<sup>2</sup>). The thickness was calculated from the mass difference of the aluminum plates between before and after the deposition. Powder samples (<7  $\mu\text{m}$  in grain size) were also prepared for gamma ray irradiation.

A <sup>60</sup>Co gamma ray source at Takasaki Research Institute of Japanese Atomic Energy Agency (JAEA) was used to irradiate the samples with a dose rate of 467.7 Gy/h to doses up to about 10 kGy.

A Tandem accelerator at Takasaki Ion Accelerators for Advanced Radiation Application (TIARA) of Japan Atomic Energy Agency (JAEA) was used to implant the  $\text{He}^+$  ions into the deposited samples where the aluminum plates were attached on a copper plate. The  $\text{He}^+$  ions with an energy of 4 MeV was implanted with a current of about 200 nA in duration of 2 to 80 seconds to an area of about 2 cm<sup>2</sup>. The number of implanted  $\text{He}^+$  ions were monitored by integrating the current during the implantation where the total number ranged  $1.2 \times 10^{12}$  to  $5.0 \times 10^{13}$  ions/cm<sup>2</sup>. The samples were peeled from an aluminum plate for ESR measurements. The doses given to the samples were calculated by dividing the total energy of  $\text{He}^+$  ions, that have passed the surface of the barite samples deposited on an aluminum plates, by the mass of barite sample on the aluminum plate, the surface of which the  $\text{He}^+$  ions passed, and in the volume of which  $\text{He}^+$  ions reached (10  $\mu\text{m}$  from the surface, see **Fig. 1**).

**Table 1.** The samples for the present study. The efficiency is the slope of the tangent at the point where signal intensity is zero on the saturating exponential curve of the dose response. The  $k$ -values are obtained as the ratio of the efficiencies to He ion implantation and to the gamma ray irradiation.

Sample No.	Site	Sampling Location		Depth (m)	Cruise No.	ESR age (year)	Efficiency to He ion dose (Gy <sup>-1</sup> )	Efficiency to gamma ray dose (Gy <sup>-1</sup> )	k-value
		Latitude (N)	Longitude (E)						
HPD#1621R07	Hatoma Knoll	24°51.33'	123°50.33'	1499	KY14-02	1480 <sup>+190</sup> <sub>-160</sub>	0.020 ± 0.001	0.37 ± 0.04	0.053 ± 0.006
HPD#1358R03	Iheya North Knoll	27°47.46'	126°53.73'	982	NT12-06	6300 <sup>+500</sup> <sub>-420</sub>	0.011 ± 0.001	0.11 ± 0.02	0.102 ± 0.022
Synthetic							0.022 ± 0.001	1.28 ± 0.07	0.017 ± 0.001



**Fig. 1.** The set-up of the experiment for  $\text{He}^+$  ion implantation showing the cross section of the barite sample deposited on an aluminum plate.

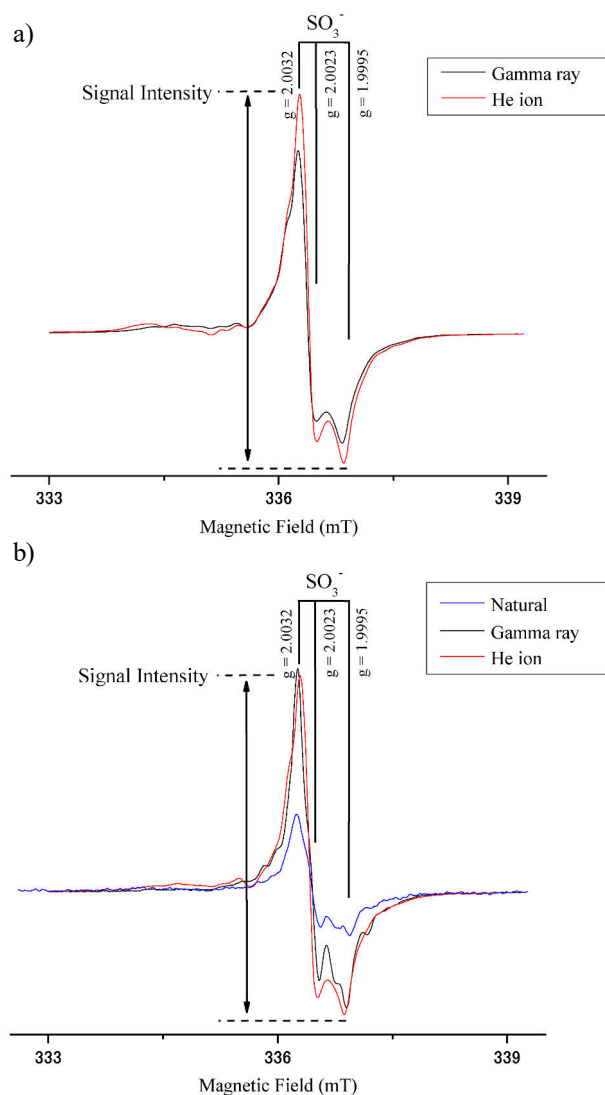
After gamma ray irradiation or  $\text{He}^+$  ion implantation, the samples were measured at room temperature with an ESR spectrometer (JES-PX2300) with a microwave power of 1 mW and the magnetic field modulation amplitude of 0.1 mT (Toyoda *et al.*, 2011).

### 3. RESULTS AND DISCUSSIONS

The ESR spectra observed in the synthetic and a natural hydrothermal (HPD#1358R03) barite samples are shown in **Fig. 2**. The principal  $g$ -values are obtained from this powder spectrum to be 1.9995, 2.0023, and 2.0031, being consistent with the  $g$ -values for  $\text{SO}_3^-$  radical obtained by Ryabov (1983), which are 1.9995, 2.0023, and 2.0032. The shapes of the signals observed in the samples irradiated by gamma rays, those in implanted by  $\text{He}^+$  ions, and that in natural samples are essentially identical.

The ESR signal intensities were enhanced both by gamma ray irradiation and by  $\text{He}^+$  ion implantation as shown in **Fig. 3** where single saturating exponential curves were fitted to the dose responses for natural samples and exponential with linear functions (Duval, 2012) were for the synthetic sample. The scattering of points for  $\text{He}^+$  ion implantation is much smaller than that in Toyoda *et al.* (2012). The error for the measurement of the signal intensity was typically within 1%, as the standard deviation of the intensities for separate aliquots of a same sample.

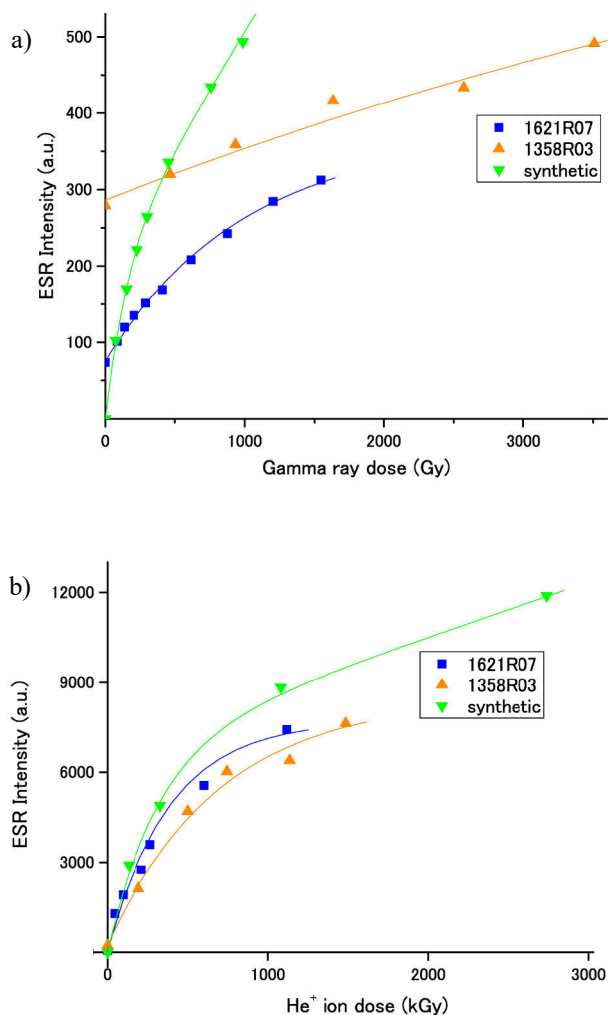
The both dose responses follow the saturating exponential curves (or plus linear), which would indicate that the  $k$  values, which are the ratio of the slopes of the tangent on the dose response curves of  $\text{He}^+$  ion implantation and of gamma ray irradiation, may change with dose. The range of the equivalent doses in the actual dating samples is quite low, typically up to about 400 y (e.g. Fujiwara *et al.*, 2015). Therefore, the slope values of the tangents at zero dose are adopted as the formation efficiencies of the signal as the following. For the  $\text{He}^+$  ion dose responses, the inherited natural doses can be neglected as the given doses are much larger (**Fig. 3**), the slope values of the tangents at given zero dose were adopted. For the gamma ray dose responses, the slope value of the tangent at given zero dose was similarly adopted for the synthetic samples. However, for the young and old natural samples, as the samples have already some given natural doses, the



**Fig. 2.** The ESR signals observed in (a) a synthetic barite sample and (b) a natural sea-floor hydrothermal barite sample (HPD#1358R03).

values at the negative equivalent dose values, at which the signal intensity is zero, were adopted as the efficiencies of signal formation by gamma rays, to be used for calculation of the  $k$ -values. The obtained efficiency values (the slopes of the tangents) are listed in **Table 1**. The  $k$  values are, then, calculated using the **Eq. 1.4** to be  $0.053 \pm 0.006$  for HPD#1621R07 (the young natural sample), to be  $0.102 \pm 0.022$  for HPD#1358R03 (the natural old sample), and to be  $0.017 \pm 0.001$  for the synthetic sample as shown in **Table 1**.

The obtained  $k$ -value is largest in the old natural sample while it is smallest in the synthetic sample. As long as we use natural samples for dating, those for the natural samples should be the candidate. The both efficiency values to  $\text{He}^+$  ion dose and to the gamma ray are smaller in the old natural sample than in the young sample as shown in **Table 1** and also in **Fig. 3**. The old sample has



**Fig. 3.** Dose responses of the ESR intensity of  $SO_3^-$  signal to (a) gamma ray and (b)  $He^+$  ion implantation doses.

already received substantial amount ( $D_E = 3240$  Gy) of natural beta and gamma and also alpha doses. Therefore, for additional gamma ray dose, it is possible that the sensitive volume in barite, where the ESR signal can be formed, would be less than the young sample, as the part of the volume damaged by the natural alpha particles would be not any more sensitive to the additional gamma ray dose. Although the same mechanism should happen for the volume with natural beta and gamma rays (the volume should be smaller), the latter effect can be considered by fitting a saturating exponential curve, while the former mechanism cannot be simulated by the gamma ray irradiation as the type of the radiation is different. The response to additional gamma ray irradiation would, therefore, be smaller because the sensitive volume, hence mass, is smaller. The same mechanism would be also for the additional  $He^+$  ion implantation. The additional  $He^+$  ion implantation will overprint the damage onto the volume with signals created by natural beta and gamma rays,

therefore, the response in the older sample would be smaller than in the young sample.

Tentatively, based on the present results, the k-value,  $0.053 \pm 0.006$ , obtained for the young natural sample should be adopted for young barite samples (typically up to 1500 years). This value is 19% higher than the previously reported value of  $0.043 \pm 0.018$  (Toyoda *et al.*, 2012). With this value, the ESR ages will typically become about 10% younger than the ones with the old k-values where currently, this difference will not affect the geological interpretations. This k-value would be much more realistic than previous value in the sense that the present value was obtained with using the natural barite extracted from a hydrothermal sulfide deposit which is actually used in dating while Toyoda *et al.* (2012) examined the barite crystal occurring on land. This value is very similar with the value, 0.052, for calcite speleothem obtained by Lyons and Brennan (1991), and in the same order with those for mollusk shells, 0.07 to 0.10 (Grün, 1985; Grün and Katzenberger, 1994) and with those for corals, 0.05 to 0.15 (Ikeya and Ohmura, 1983; Radtke *et al.*, 1988; Grün *et al.*, 1992; Malmberg and Radtke, 2000).

However, further works would still be needed, firstly, to confirm that the k-value for the present young natural sample, HPD#1621R07, is reproducible in other natural samples. The efficiency value to  $He^+$  ion dose in the synthetic sample is consistent with the one in the young natural sample as shown in Table 1 while the value to gamma ray dose is higher in the synthetic than in the natural young sample, making the k-value lower. This efficiency value, the saturated intensity value, and the dose value at which saturation starts (typically indicated by  $D_0$ ) in the dose response would directly be related with the formation mechanism. Indicated by the responses to gamma rays observed in the present study (Fig. 3), it is not so simple as other signal such as one observed in a clay mineral where the higher the saturation value corresponds to the lower the efficiency (Allard and Muller, 1998). In barite, the efficiency to  $He^+$  ion, hence, alpha dose, is constant while the one to beta and gamma might be sample dependent, which would possibly be due to chemical and physical conditions of barite formation. Secondary, as indicated by the results for the old natural sample, the k-value possibly changes with the alpha dose. It is necessary to examine other samples with variety of ages and also to examine the dose response to the gamma rays in  $He^+$  ion implanted samples in order to investigate such dependence. If this is the case, some complex numerical calculation with iteration with changing k-value may be necessary.

#### 4. SUMMARY

The value of alpha effectiveness was investigated by the experiments with 4 MeV  $He^+$  ion implantation and gamma ray irradiation to three barite samples. The ob-

tained dose response of the  $\text{SO}_3^-$  signal to the  $\text{He}^+$  ion dose is much better than the one previously reported by Toyoda *et al.* (2012). Tentatively, a value of  $0.053 \pm 0.006$ , which was obtained for a young natural barite sample, should be adopted as the k-value. However, the present study also found that the value is sample dependent possibly on the age, therefore, the value should be used for young samples up to 1500 years. Further studies on the k value are still necessary.

## ACKNOWLEDGMENTS

This work has been supported by the Inter-University Program for the Joint Use of JAEA Facilities, and also partly by MEXT-Supported Program for the Strategic Research Foundation at Private Universities (2011-2015), and  $\text{He}^+$  ion implantation experiments were supported by the Inter-University Program for the Joint Use of the JAEA (Takasaki), Grant No. 15016 to H.N.

## REFERENCES

- Allard T and Muller J-P, 1998. Kaolinite as an in situ dosimeter for past radionuclide migration at the Earth's surface. *Applied Geochemistry* 13(6): 751–765, DOI 10.1016/S0883-2927(98)00011-0.
- Duval M, 2012. Dose response curve of the ESR signal of the Aluminum center in quartz grains extracted from sediment. *Ancient TL* 30: 41–49.
- Fujiwara T, Toyoda S, Uchida A, Ishibashi J, Nakai S and Takamasa A, 2015. *ESR dating of barite in sea-floor hydrothermal sulfide deposits in the Okinawa Trough*. In Ishibashi, J., Okino, K. and Sunamura, M. eds., *Subseafloor Biosphere Linked to Global Hydrothermal Systems*; TAIGA Concept, Springer, Tokyo, 369–386.
- Grün R, 1985. Beiträge zur ESR-Datierung (Contributions to ESR dating). *Sonderveröffentlichungen des Geologischen Instituts der Universität zu Köln* 59: 1–157. In German.
- Grün R and Katzenberger-Apel O, 1994. An alpha irradiator for ESR dating. *Ancient TL* 12: 35–38.
- Grün R, Radtke U and Omura A, 1992. ESR and U-series analyses on corals from Huon Peninsula, New Guinea. *Quaternary Science Reviews* 11: 197–202, DOI 10.1016/0277-3791(92)90063-E.
- Ikeya M and Ohmura K, 1983. Comparison of ESR ages of corals from marine terraces with  $^{14}\text{C}$  and  $^{230}\text{Th}/^{234}\text{U}$  ages. *Earth and Planetary Science Letters* 65: 34–38, DOI 10.1016/0012-821X(83)90187-5.
- Kasuya M, Kato M and Ikeya M, 1991. *ESR signals of natural barite ( $\text{BaSO}_4$ ) crystals: possible application to geochronology*. In *Essay in Geology, Prof. Nakagawa Commemorative Volume*: 95–98.
- Lyons RG and Brennan GJ, 1991. Alpha/gamma effectiveness ratios of calcite speleothems. *Nuclear Tracks and Radiations Measurements* 18(1/2): 223–227, DOI 10.1016/1359-0189(91)90116-Y.
- Malmberg R and Radtke U, 2000. The alpha-efficiency of corals and its importance for ESR-dating. *Radiation Measurements* 32: 747–750, DOI 10.1016/S1350-4487(00)00079-2.
- Okumura T, Toyoda S, Sato F, Uchida A, Ishibashi J and Nakai S, 2010. ESR Dating of marine barites in chimneys deposited from hydrothermal vents. *Geochronometria* 37: 57–61, DOI 10.2478/v10003-010-0019-z.
- Radtke U, Grün R and Schwarcz HP, 1988. Electron spin resonance dating of the Pleistocene coral reef tracts of Barbados. *Quaternary Research* 29: 197–215, DOI 10.1016/0033-5894(88)90030-0.
- Ryabov ID, Bershov LV, Speranskiy AV and Ganeev IG, 1983. Electron paramagnetic resonance of  $\text{PO}_3^{2-}$  and  $\text{SO}_3^-$  radicals in anhydrite, celestite and barite: the hyperfine structure and dynamics. *Physics and Chemistry of Minerals* 10: 21–26, DOI 10.1007/BF01204322.
- Takamasa A, Nakai S, Sato F, Toyoda S, Banerjee D and Ishibashi J, 2013. U-Th radioactive disequilibrium and ESR dating of a barite-containing sulfide crust from South Mariana Trough. *Quaternary Geochronology* 15: 38–46, DOI 10.1016/j.quageo.2012.12.002.
- Toyoda S, Sato F, Banerjee D and Ishibashi J, 2011. Characteristics of the Radiation Induced ESR Signals in Barite. *Advances in ESR applications* 27: 4–6.
- Toyoda S, Sato F, Nishido H, Kayama M and Ishibashi J, 2012. The alpha effectiveness of the dating ESR signal in barite. *Radiation Measurements* 47: 900–902, DOI 10.1016/j.radmeas.2012.04.016.
- Toyoda S, Fujiwara T, Uchida A, Ishibashi J, Nakai S and Takamasa A, 2014. ESR dating of barite in sulphide deposits formed by the sea-floor hydrothermal activities. *Radiation Protection Dosimetry* 159: 203–211, DOI 10.1093/rpd/ncu136.